

Modelling the hydrological response of a small Mediterranean forested catchment: exploring the potential influence of the riparian-stream connection

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Mediterranean catchments

- Characteristics:
 - Share hydrological processes from both *wet* and *dry* environments
 - Large range of weather conditions that lead to a complex stream hydrology
 - ✓ *High variability* in the annual water balance
 - ✓ Seasonal pattern in hydrological behaviour:



• Long summer <i>dry period</i> \rightarrow	Switching behaviour of the permanent saturated zone. (<i>Gallart et al., 2002; Marc et al., 2001; Piñol et al., 1997</i>)
• Wetting-up period \longrightarrow	Appearence of a perched water table quite fast draining.(Burch et al.1987; Ocampo et al., 2006; Taha et al., 1997)
 Autumn/winter wet period —> 	Recharge of the deeper saturated areas. (Butterworth et al., 1999; Pilgrim et al., 1997)

Concepts and ideas developed by modellers for humid climate usually fail when applied to semi-arid regions and lead in many cases to unsatisfactory results

(Bernal et al., 2004; Bonell, 1993; Latron et al.; 2003, Pilgrim et al., 1988)

The challange of this study was to improve the representation and understanding of flow processes in Mediterranean catchments with special attention to the transition period

✓ Hydrological management of this critical areas

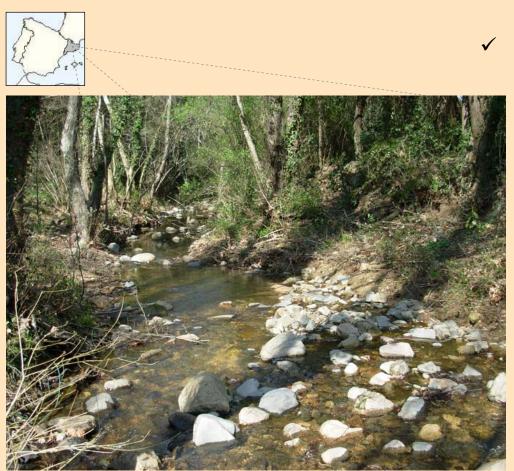
(Chiew et al., 2002)

✓ Good prediction of geochemical and ecological responses

(Schlesinger et al., 2006)

Study Site

The Fuirosos catchment (13 km²):



- ✓ is located in the North-East of Spain (latitude 41°42'N, longitude 2°34').
 - is an *almost pristine*, undisturbed forested catchment, which drains an intermittent stream. There is little agricultural activity and no urban areas.

Within the catchment there are *four small reservoirs* for human and cattle water supply.

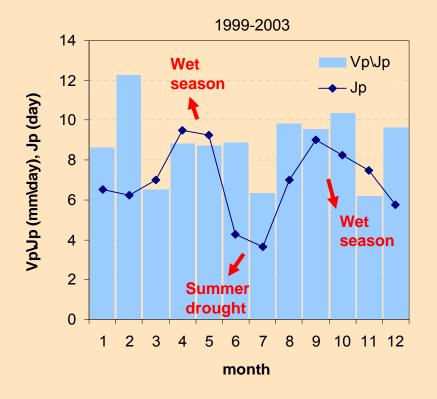
The main rock type is *leucogranite* (50.9%) followed by *granodiorite* (21.1%) and *sericitic schists* (23.5%). At the valley bottom there is an identifiable *alluvial zone*, with a well developed *riparian zone*.

The *forest* covers the 90% of the total cacthment area.

Study Site

General water balance analysis:

- ✓ Monthly mean temperature ranges from 3°C in January to 24°C in August (Bernal et al., 2004)
- ✓ Average annual precipitation is 750 mm (*Ninyerola et al., 2000*), and the annual average number of rainy days is 81 (P≥4 mm)



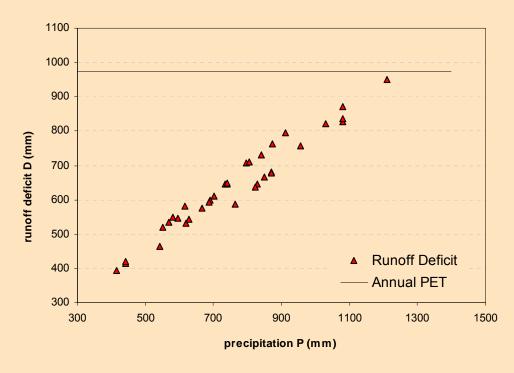
Jp : monthly number of rainy days Vp\Jp : total precipitation per rainy days

There are two identifiable *wet periods*: one during *spring* and the other during *autumn;* and a *summer drought*.

Study Site

General water balance analysis:

- Average annual potential evapotranspiration (PET) is approximately 975 mm, according to the Penman Method
- The average annual runoff deficit (D) is approximately 640 mm, with a Q/P coefficient of 15%



NO REASONS TO CONSIDER

GROUNDWATER OUTFLOW!!

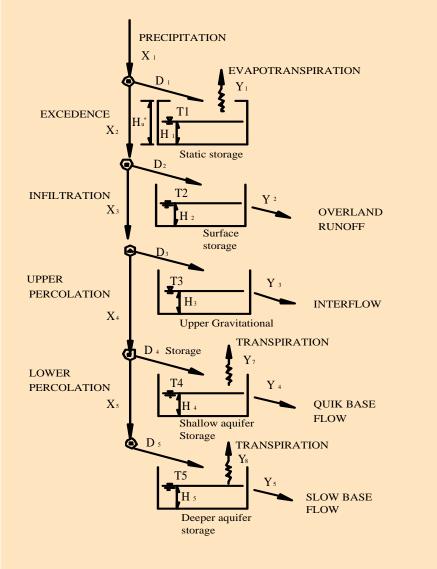
The graph points out that the *runoff deficit* can be basically related with *evapotranspiration*, since the annual precipitation is lower than the evaporative demand.

Water supply in the wettest year is still not enough to satisfay the PET.

Bedrock characteristics also support this hypothesis

Model descripion

SD4 : Semidistributed 4-response Model



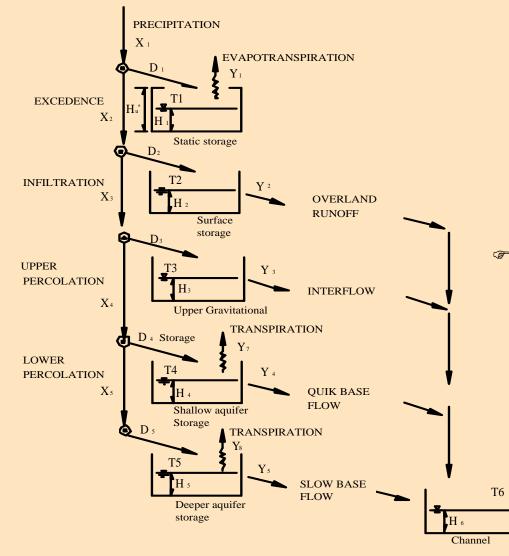
1. Static tank: Initial abstractions and water retained by capillary force

$$D_{1}(t) = \min \begin{cases} X_{1}(t) \cdot \left(1 - \frac{H_{1}(t)}{H_{u}^{*}}\right)^{2}; H_{u}^{*} - H_{1}(t) \\ Y_{i=2,...,5}(t) = \\ \end{cases}$$
2. Surface tank: Water_{1,flowing over the right of the the tank of tank of the tank of tank of

- 3. Gravitational tank: Water flowing into a soil-gravel layer, horizon A, as interflow.
- 4. Shallow aquifer: This tank represents a perched aquifer that may appear in the upper weathered bedrock layer, *horizon B*. The flow that is released from it, is thought to be a key process during the wetting-up period.

Model descripion

SD4 : Semidistributed 4-response Model



5. Deep aquifer: Represents the permanent saturated zone into a deeper weathered bedrock layer. It is thought to be constitued by several bedrock depressions which may exert a significant control on water mobility.

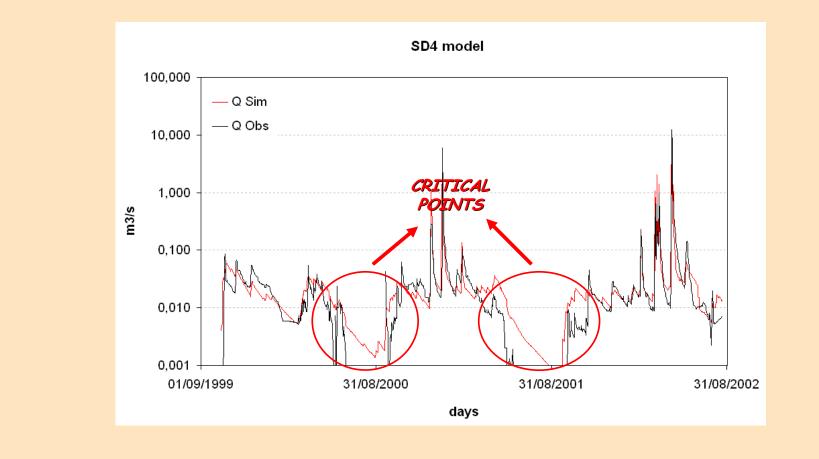
$$Y_{i=2,...,5}(t) = \frac{H_{i=2,...,5}(t)}{t_{i=1,...,4}}$$

Both aquifers are thought to be accesible by the root system.

Conceptual scheme for each HRU

 ✓ All the tanks are described as linear storages and drain directly to the stream tank

Results with the SD4 Model



Nash Index:0.77SimTotal balance Vol. Err. :4.6%ObsN. of days with Sim. Q < $0,001 \text{ m}^3/\text{s}$:28N. o

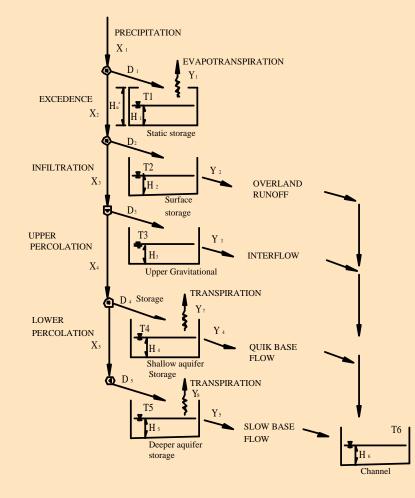
 Sim. Peak:
 8.7 m³/s

 Obs. Peak:
 10.9 m³/s

 N. of days with Obs. Q < 0,001 m3/s :</td>
 220

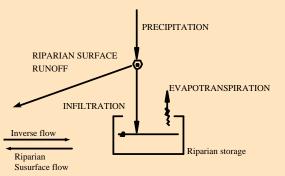
Model description

SD4-R : Semidistributed 4-response Model plus a riparian tank

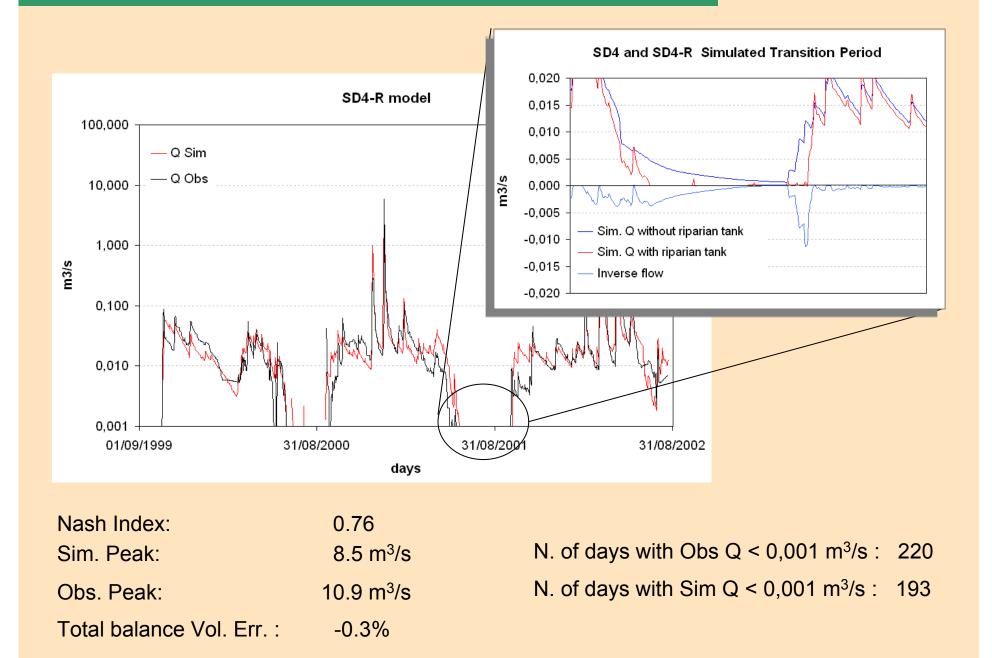


✓ Exchanges of water are generated accordign to the difference between the river stage (d) and the riparian gruondwater head (e), the hydraulic conductivity and the effective porosity of the soil.

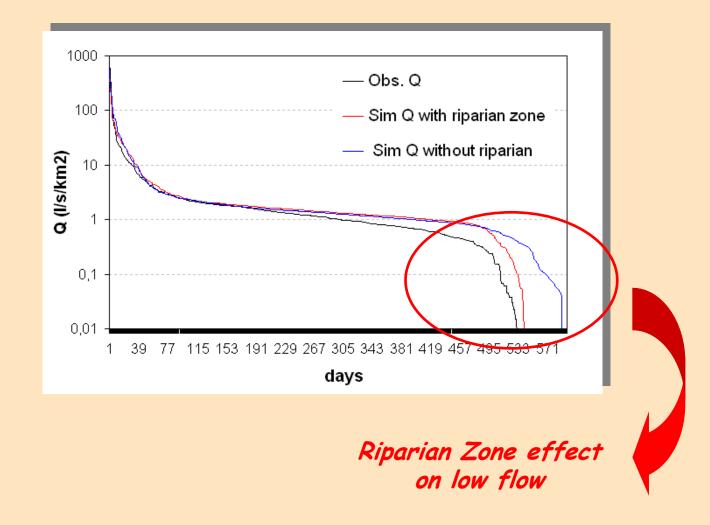
$$F_{d,i} = \pm K_r \cdot \left(\frac{e-d}{m}\right) \cdot \left(2 \cdot f \cdot c\right)$$



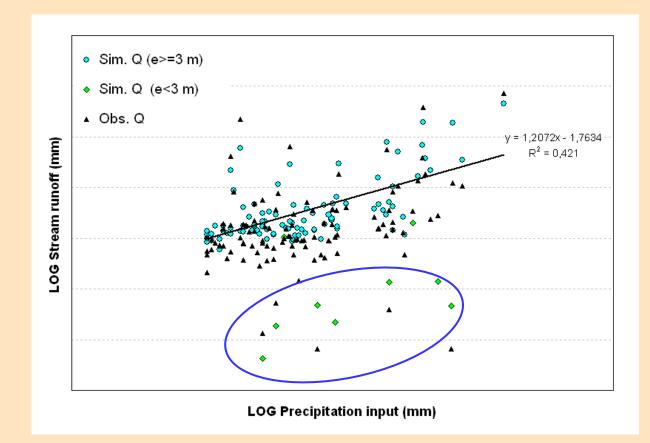
Results with the SD4-R Model



Results Analysis



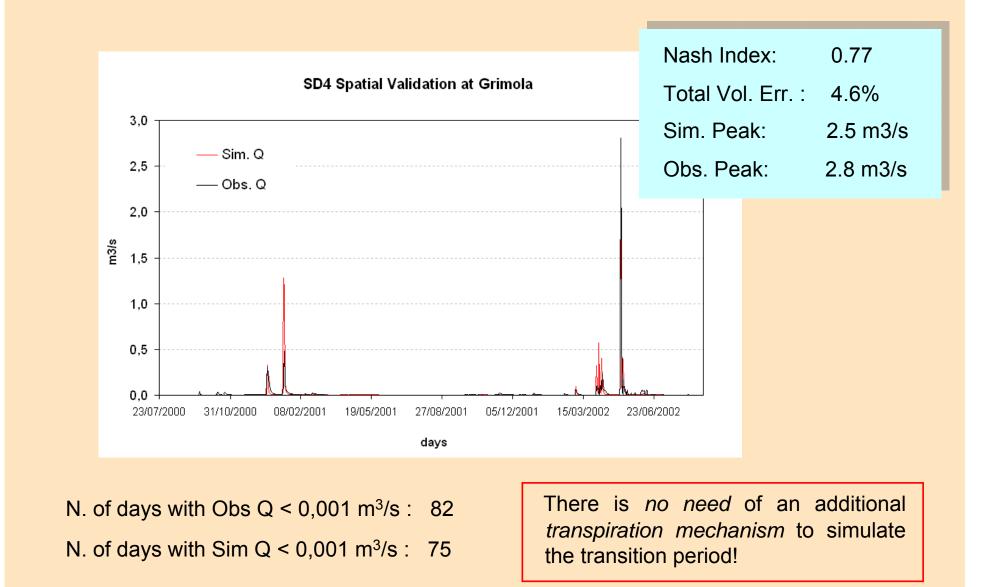
Results Analysis



Water availability in the riparian area affects the relationship between rainfall inputs and stream runoff.

Stream runoff and the rainfall input get better correlated only after riparian groundwater store raises its water table level above the streambed

Spatial Validation



✓ Temporal validation had similar performance

Models developed for humid climate can not capture the characteristic *inter-annual* and *intra-annual variability* of Mediterranean catchments.

The transition period from dry to wet condition has been noted as a critical point to be reproduce by available rainfall-runoff models.

In the analysis of an *intermittent stream* the riparian zone may represent an *important mechanism* to explain the transition period.



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